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## Recent changes in chironomid communities and hypolimnetic oxygen conditions relate to organic carbon in subarctic ecotonal lakes

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1    **Recent changes in chironomid communities and hypolimnetic oxygen conditions**  
2    **relate to organic carbon in subarctic ecotonal lakes**

3

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26   **Abstract**

27

28   A key question in aquatic elemental cycling is related to the influence of bottom water oxygen  
29   conditions in regulating the burial and release of carbon under climate warming. In this study, we  
30   used head capsules of Chironomidae larvae to assess community and diversity change between the  
31   past (estimated as Pre-Industrial Period) and present and to reconstruct changes in hypolimnetic  
32   oxygen conditions from 30 subarctic ecotonal lakes (northeastern Lapland) using the top-bottom  
33   paleolimnological approach applying surface sediment (topmost 0-2 cm) and reference (4-5 cm)  
34   samples. Subsequently, we tested the findings against dissolved organic carbon (DOC)  
35   concentration of the sites. We found that the benthic communities were statistically dissimilar  
36   between the past and the present with largest changes occurring in the more transparent oligo-  
37   mesohumic lakes. However, murky polyhumic lakes displayed uniformly a decrease in diversity.  
38   The chironomid-inferred oxygen values showed a general decrease towards the present with largest  
39   shifts in low-DOC lakes, whereas no significant changes were found in the hypolimnetic oxygen  
40   conditions of high-DOC lakes, which were often located in wetland areas. These finding suggest  
41   that lakes associated with constant organic carbon inputs are more resilient toward climate-induced  
42   reductions in hypolimnetic oxygen.

43

44   *Keywords:* Chironomidae; dissolved oxygen; Lapland; DOC; paleolimnology; top-bottom approach

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## 51 **1 Introduction**

52

53 Carbon storage in high-latitude peatlands represents a major share of the global soil carbon pool  
54 (Tranvik et al., 2009; Schuur et al., 2015). With permafrost thawing and subsequent formation of  
55 thaw lakes, larger emissions of carbon dioxide (CO<sub>2</sub>) and methane (CH<sub>4</sub>) are expected (Walter et  
56 al., 2006; Wauthy et al., 2018). Since lakes are important sources of greenhouse gases (Aben et al.,  
57 2017), in the context of climate change, it is essential to understand how carbon release is regulated  
58 by different levels of oxygen availability at the surface-sediment interface of lakes (Liikanen et al.,  
59 2002). Although there is CH<sub>4</sub> formation also in oxic water layers (Tang et al., 2016), CH<sub>4</sub> is mostly  
60 produced by anoxic decomposition of organic carbon and, importantly for the ongoing climate  
61 change, it has a multifold atmospheric warming potential compared to CO<sub>2</sub> (Deemer et al., 2016).

62 While in lakes with well-oxygenated bottom waters a significant part of the produced  
63 CH<sub>4</sub> is oxidized into CO<sub>2</sub> and not emitted to the atmosphere, a contrasting situation occurs in  
64 oxygen depleted lakes. In addition to hypolimnetic oxygen, there appears to be a consistent  
65 temperature dependency of CH<sub>4</sub> fluxes across microbial to ecosystem scales (Yvon-Durocher et al.,  
66 2014) that well-represents the cascading climate change impacts and feedback systems. In  
67 particular, methanogenesis, carried out by strictly anaerobic Archaea, is the terminal step in the  
68 remineralization of organic matter and, like most other forms of metabolism, methanogenesis is  
69 temperature-dependent (Yvon-Durocher et al., 2014). Therefore, since climate warming is  
70 emphasized in subarctic and arctic regions (Linderholm et al., 2018), northern lakes play a  
71 significant role in the ongoing climate change through their functions in carbon release and  
72 sequestration.

73 Chironomid (Insecta: Diptera: Chironomidae) larvae take part in benthic processes  
74 that are essential for lake ecosystem functions (Benke and Huryn, 2010; Hölker et al., 2015). Most  
75 importantly, chironomids living in the water-sediment interface have a major role in the

76 biogeochemical cycling of carbon, nitrogen and phosphorus (Nogaro et al., 2008; Belle et al., 2017).  
77 The functional roles of chironomids are mostly reflected by their feeding preferences (Heino, 2008;  
78 Luoto and Nevalainen, 2015), but their taxonomical compositions typically respond to climate  
79 conditions at the regional scale (Brooks, 2006) and hypolimnetic oxygen and nutrient conditions at  
80 the local/site-specific scale (Brodersen and Quinlan, 2006). In Finnish Lapland, the most important  
81 environmental factors controlling chironomid distribution and abundance are temperature, organic  
82 content of the sediment, nutrients/oxygen and water depth (Olander et al., 1997; Nyman et al., 2005;  
83 Luoto et al., 2016). Hence, chironomids contribute as sensitive indicators of climate changes,  
84 limnology and elemental cycling. The influence of temperature on chironomids can be direct  
85 (metabolism) or indirect (e.g. habitat) and mediated through air (flying adults) or water (aquatic  
86 larval stage) temperature (Eggermont and Heiri, 2012), whereas the influence of hypolimnetic  
87 oxygen is related to the species-specific respiration rates and oxygen demands (Brodersen et al.,  
88 2008). Using the transfer function approach (calibration-in-space), it is possible to provide long-  
89 term quantitative reconstructions of these variables (Quinlan and Smol, 2002; Luoto et al., 2017;  
90 Wohlfarth et al., 2018).

91           In this study, we investigate chironomid assemblages and changes in chironomid-  
92 inferred hypolimnetic oxygen conditions in 30 subarctic lakes in northern Finnish Lapland using the  
93 top-bottom (or before-after) paleolimnological approach (Quinlan and Smol, 2002), where the  
94 surface sediment samples (top) represent the present and the reference samples (bottom) the past.  
95 We examine the findings against the modern dissolved organic carbon (DOC) gradient of the lakes  
96 to find out potential differences caused by catchment characteristics, with special interest on the  
97 influence of peatlands.

98

## 99 **2 Material and Methods**

100

## 101 2.1 Study sites and sampling strategy

102

103 The study sites include 30 lakes from northeastern Finnish Lapland (68°47'–69°55'N) with a  
104 catchment gradient from mixed pine and birch forest to mountain birch woodland and barren tundra  
105 (Fig. 1). The study area is located in a sporadic permafrost region. All the lakes are small and  
106 shallow (<7.5 m) and distributed along a truncated mean July air temperature transect from 12.3 to  
107 13.3 °C. While minimizing the temperature gradient, the lakes were sampled to characterize  
108 different catchment types from forest and bog environments to tundra vegetation to encompass a  
109 broad dissolved organic carbon (DOC) gradient from 1.7 to 16.6 mg l<sup>-1</sup>. The sampling strategy was  
110 originally designed to study spatio-temporal zooplankton responses to UV radiation (the DOC  
111 screening effect) (Nevalainen et al., 2018) and to investigate the sources and controls of organic  
112 carbon in these lakes (Rantala et al., 2016a). The dataset consists of 15 oligohumic (DOC <4 mg l<sup>-1</sup>  
113 <sup>1</sup>), 10 mesohumic (DOC = 4-6 mg l<sup>-1</sup>) and 5 polyhumic (DOC >6 mg l<sup>-1</sup>) lakes. Total phosphorus in  
114 the lakes varies from 5 to 24 µg l<sup>-1</sup>, total nitrogen from 138 to 806 µg l<sup>-1</sup> and pH from 5.1 to 8.4.  
115 Environmental characteristics and limnology of the study sites are described in detail in Rantala et  
116 al. (2016a). The lakes and their codes (numbers) are the same as in Rantala et al. (2016a), with the  
117 exception that one tundra site (#34) was removed from this study due to low number of chironomid  
118 head capsules.

119 Surface (top, 0–2 cm in core depth) and reference (bottom, 4–5 cm) sediment samples  
120 were collected from the centres of the lakes using a Limnos gravity corer (Kansanen et al., 1991) in  
121 July 2014. According to several sediment cores dated from the study area (Luoto & Sarmaja-  
122 Korjonen, 2011; Luoto et al., 2017b; Rantala et al., 2016b), the average sediment accumulation in  
123 the lakes refer to past 20 years in the “top” surface sediment samples and to 150 years in the  
124 “bottom” reference samples. The top and bottom samples are discussed herein as modern and Pre-  
125 Industrial age samples, respectively, but since the samples were not independently dated, we

126 emphasize that these general timeframes are only tentative and the bottom sample may not always  
127 represent Pre-Industrial times. Although the widely employed top-bottom (or before-and-after)  
128 approach allows for efficient spatio-temporal regional assessment of environmental change, the  
129 disadvantage of this approach is that there are no continuous data on the timing or rate of changes  
130 occurring between the two points in time represented by those samples (Smol, 2017). In addition, it  
131 cannot be assumed that there is only one Pre-Industrial/reference environmental state from which  
132 lakes have deviated (Perren et al., 2009). These limitations of the applied approach should be kept  
133 in mind when interpreting the data.

134

## 135 *2.2 Chironomid analysis*

136

137 Standard methods were applied in fossil chironomid analysis (Brooks et al., 2007). The wet  
138 sediment was sieved through a mesh (100- $\mu$ m) and the residue was examined under a  
139 stereomicroscope. Larval head capsules were extracted and mounted permanently with Euparal on  
140 microscope slides. Taxonomic identification following Brooks et al. (2007) was performed under a  
141 light microscope. The minimum chironomid head capsule number per sample was set to 50 (Heiri  
142 and Lotter, 2001; Larocque, 2001; Quinlan and Smol, 2001). Two split head capsules were  
143 considered as one individual. The surface sediment chironomid assemblages have been previously  
144 published in Luoto et al. (2016), whereas the reference samples were analyzed in this study by the  
145 same chironomid analyst using identical methodology and taxonomy.

146

## 147 *2.3 Numerical methods*

148

149 Bray-Curtis similarity was used as a measure to assess the difference between present and past  
150 chironomid communities. In this measure, 0 refers to a completely different community and 1 to an

151 analogous community. N2 (Hill, 1973) was used as a diversity measure, corresponding to effective  
152 number of occurrences, i.e. the effective diversity of a sample's community.

153 Minimum hypolimnetic dissolved oxygen (DO) was reconstructed using a 30-lake  
154 chironomid-based calibration model for Finland, where the calibration sites range from anoxic  
155 ( $O_2 < 0.5 \text{ mg l}^{-1}$ ) to hypersaturated sites ( $O_2 = 18.1 \text{ mg l}^{-1}$ ) (Luoto and Nevalainen, 2011; Luoto and  
156 Salonen, 2010). The weighted averaging partial least squares model has a cross-validated (leave-  
157 one-out) coefficient of determination of 0.74 and a root mean squared error of prediction of  $2.3 \text{ mg}$   
158  $\text{l}^{-1}$ . The suitability of the calibration model for the chironomid assemblages in the surface and  
159 bottom samples were assessed using representation of fossil taxa in the calibration set (cutoff for  
160 good representativeness 95%) and the modern analogue technique (MAT) using squared chi-square  
161 distances of the 10 closest modern analogues (cutoff for a good modern analogue 5% squared chi-  
162 square distance) (Birks et al., 2014). Furthermore, Generalized Linear Modeling (GLM) was used to  
163 assess taxa that significantly ( $P \leq 0.05$ ) respond to oxygen conditions. The GLMs were run using  
164 Poisson distribution. Samples where majority of the taxa have significant relationship with the  
165 environmental variable of interest can be considered to produce reliable reconstructions (Rees et al.,  
166 2008; Luoto et al., 2010).

167

### 168 **3 Results**

169

170 The counting sum of chironomid individuals varied between 50 and 60 in the sediment samples.  
171 The most common and abundant chironomids in the bottom samples were the same as in the surface  
172 sediment samples (Fig. 2), including *Psectrocladius septentrionalis*-type (bottom = 27 occurrences,  
173 mean abundance 12.6%, maximum abundance 48.9%; top = 23, 12.0%, 45.2%), *P. sordidellus*-type  
174 (bottom = 28, 12.4%, 39.3%; top = 27, 10.2%, 23.6%) and *Tanytarsus lugens*-type (bottom = 28,  
175 8.0%, 28.3%; top = 25, 5.9%, 19.0%). The most common taxa with largest decreases in their mean



176 abundance from bottom to top samples included *Polypedilum nubeculosum*-type (-2.5%), *P.*  
177 *sordidellus*-type (-2.2%) and *T. lugens*-type (-2.1%), whereas the largest increases were observed  
178 with *Paratanytarsus penicillatus*-type (3.2%), *Microtendipes pedellus*-type (2.1%) and *Tanytarsus*  
179 *glabrescens*-type (2.0%). *Protanypus*, which was present in 11 bottom samples, was encountered  
180 only in a single surface sample. On the contrary, *Tantarsus mendax*-type that was absent in the  
181 bottom samples appeared in 10 surface samples.

182           According to the similarity analysis, oligotrophic tundra lakes had the largest shifts in  
183 taxa assemblages between the past and present (average Bray-Curtis similarity 0.44), whereas  
184 polyhumic lakes, located in pine and birch forest (0.50) and mountain birch woodland (0.55),  
185 exhibited less changes (Table 1; Fig. 3). However, unlike oligohumic and mesohumic sites, the  
186 polyhumic sites showed consistently decreased diversity (Table 1; Fig. 3).

187           The chironomid-inferred hypolimnetic oxygen varied between 3.0 and 13.5 mg l<sup>-1</sup> in  
188 the surface samples and between 6.6 and 17.1 mg l<sup>-1</sup> in the bottom samples (Fig. 3). In all, the  
189 average chironomid-inferred oxygen for the surface sediments was 9.4 mg l<sup>-1</sup> and for the bottom  
190 samples 12.6 mg l<sup>-1</sup>. According to the MAT, all the surface and bottom samples had good modern  
191 analogues (<5% squared chi-square distance) in the calibration set. However, the surface sample  
192 from lake #27 had taxa representativeness slightly less than the cutoff value of 95% due to absence  
193 of *Paratendipes nudisquama*-type from the calibration set. Nonetheless, since 94.7% of the fossil  
194 taxa were present in the calibration set, this did not produce a significant error source for the  
195 reconstruction. In the surface sediment samples, 71.7-100% (average 91.0%) of the taxa present  
196 (relative community abundance) had significant relationship with oxygen conditions, whereas in the  
197 reference samples 73.1-100% (average 89.3%) were significantly associated with oxygen. Hence all  
198 samples had the needed coverage of statistically significant oxygen indicators. The most abundant  
199 taxon without significant relationship with hypolimnetic oxygen was *Procladius*, which occurred in  
200 the samples with a relatively low mean abundance of 2.4% (maximum abundance 7.4%). When

201 comparing the shifts in inferred oxygen values and their relative change (Fig. 4), the polyhumic  
202 sites showed no significant changes, mesohumic sites showed decrease with only one exception (an  
203 outlier lake with a pH of 5.1) and the oligohumic sites illustrated a coherent decrease in oxygen.  
204

## 205 **4 Discussion**

### 207 *4.1 Ecological change*

209 The chironomid assemblages in the 30 study sites in northern Lapland showed clear shifts when  
210 comparing the reference (tentative age estimate 150 years in average) and modern surface sediment  
211 samples. *Paratanytarsus penicillatus*-type, which was rare in the bottom samples, showed increase  
212 especially in the low-DOC lakes (Fig. 2). Also several other collector-filterers consistently  
213 increased, such as *Tanytarsus mendax*-type and *T. glabrescens*-type, though remaining relatively  
214 low in their abundances. When interpreting the community changes, it should be noted, according  
215 to the species-area relationship (Lomolino, 2000), that the time-intervals between the top and  
216 bottom samples may differ and cause uncertainties in comparisons of rare taxa occurrences.  
217 Nonetheless, the recent increase of *T. mendax*-type is very typical in the northern lakes in the area  
218 (Shala et al., 2014). *T. mendax*-type has a high modern temperature optimum in Lapland (Luoto et  
219 al., 2014) and elsewhere in Europe (Heiri et al., 2011). Therefore, this taxon appears to represent an  
220 example of a chironomid that clearly benefits from climate warming. In contrast, the cold-indicating  
221 *Protonyctus* has disappeared from the surface sediments of the study sites suggesting that the  
222 temperature range of the lakes is no longer suitable for it to survive. In addition to temperature, *T.*  
223 *mendax*-type and *Protonyctus* have strict preferences for meso-eutrophic and oligotrophic lakes,  
224 respectively (Luoto, 2011; Medeiros and Quinlan, 2011). For chironomids, the influence of nutrient  
225 conditions is in most part mediated through the direct physiological effects of hypolimnetic oxygen

226 conditions (Brodersen and Quinlan, 2006), as oligotrophic lakes tend to be well-oxygenized and  
227 eutrophic lakes often oxygen depleted.

228           According to the Bray-Curtis similarity, the polyhumic study sites with DOC >9  $\mu\text{g l}^{-1}$   
229 had smallest and the oligo-mesohumic lakes largest community changes (Fig. 3). While displaying  
230 relatively little change in community assemblages, the polyhumic sites appeared to have uniformly  
231 suffered a loss in their diversity. Hence, the chironomid communities of the polyhumic sites have  
232 not significantly changed in their dominating taxa, instead there have been general decreases in the  
233 less abundant taxa (Fig. 2). As all the current lakes are situated in an ecotonal area, the ecological  
234 impacts of climate change are typically magnified (Pienitz and Vincent, 2000). Therefore, the larger  
235 community shifts occurring in the low-DOC lakes (Fig. 3) are not surprising owing to oligohumic  
236 and oligotrophic lakes being generally poorly buffered against environmental perturbations (Gunn  
237 et al., 2001). Overall, the findings from the chironomid assemblages suggest that ecological changes  
238 in taxonomic composition and/or diversity occur between the reference samples and the present,  
239 with direction and magnitude depending on the lake type. As the status of aquatic ecosystems as  
240 carbon sinks or sources is very likely to change owing to the ongoing climate change, the  
241 biodiversity of freshwater ecosystems across most of the Arctic will also likely be altered (Wrona et  
242 al., 2006). Accordingly, our results provide further evidence that these changes are proceeding in  
243 subarctic lakes of Lapland, where chironomids play a significant functional role through regulation  
244 of organic matter decomposition, and subsequently, take a major part in the biogeochemical cycling  
245 of nutrients (including carbon).

246

#### 247 *4.2 Oxygen change*

248

249 As the chironomid assemblages had good modern analogues in the calibration set according to the  
250 MAT and based on the GLMs the taxonomic compositions included sensitive oxygen indicators

(Luoto and Salonen, 2010), such as the low-oxygen *Sergentia coracina*-type and high-oxygen *Tanytarsus lugens*-type and *Psectrocladius* taxa, we were able to reliably reconstruct the hypolimnetic oxygen conditions from the top and bottom samples of the study sites. The results showed differences in chironomid-inferred oxygen levels between different lake types in past and modern samples (Fig. 3). Whereas the oxygen decrease from the estimated Pre-Industrial Age was 3.2 mg l<sup>-1</sup> on average, which clearly exceed the model's prediction error, largest reductions were observed in the oligohumic lakes (Figs 3-4). With the outlier (lake #5 with anomalous pH) excluded, also mesohumic lakes distinctly suffered from decreased oxygen levels. On contrast, the polyhumic lakes with present-day DOC >6 mg l<sup>-1</sup> had no or little change in their oxygen levels between the past and the present. Therefore, the larger oxygen change in the oligo-mesohumic lakes compared to polyhumic lakes is probably due to their generally poorer buffering ability against environmental changes (Gunn et al., 2001). The lower buffering ability, or resilience, of low-DOC lakes most probably relates to their light environment (more transparent water column), trophic status (less nutrients), variable limnological conditions (such as in pH) and diverse catchment characteristics compared to the murky, low-pH polyhumic lakes with boggy catchments (Korhola et al., 2002; Rautio et al., 2011; Rantala et al., 2016a).

When observing the oxygen change since the estimated Pre-Industrial Age to the present (Fig. 4) in the light of climate warming, it becomes apparent that polyhumic lakes are less influenced. This is probably due to the fact that polyhumic sites have generally been subjected to low oxygen conditions through their polyhumic lake state, especially during the winter season when oxygen is consumed under ice following the excessive decay of plant material derived from the catchment wetlands (Mathias and Barica, 1980). It has been shown that while colored lakes are typically naturally oxygen depleted, the profundal oxygen values have a weak relationship with lake trophic state (Crisman et al., 1998). Instead, the oxygen conditions in lakes are closely connected with temperature. For example, oxygen conditions were generally improved during the cold Little

Ice Age (~1300-1900 CE) but deteriorated during the warm Medieval Climate Anomaly (800-1300 CE) and also under the present climate warming in oligohumic and polyhumic boreal lakes of Finland (Luoto and Salonen, 2010). The current results show that even though the hypolimnetic oxygen levels have decreased in the oligo-mesohumic lakes, none of the lakes have become anoxic thus far (Fig. 3). However, despite the fact that the bottom reference samples represent a time period that may correspond to the Little Ice Age and may consequently represent naturally higher values, the decreasing oxygen trend in these lakes is evident, and the future therefore holds a potential risk that the lakes may become oxygen depleted as climate warming proceeds.

284

## 285 **5 Conclusions**

286

The chironomid assemblages in the past and modern sediment samples of the 30 lakes showed dissimilar communities between the past and the present with largest shifts in the more transparent oligo-mesohumic lakes. Although there was no significant change in the taxonomic communities of the murky polyhumic lakes, they displayed a decrease in diversity driven by disappearance of rare specialized taxa. The chironomid-inferred hypolimnetic oxygen values generally decreased from the Pre-Industrial Age to the present, with largest change in low-DOC lakes but no significant changes were found in the oxygen conditions of high-DOC lakes. Since the polyhumic study lakes were often located in wetland areas, our findings suggest mechanisms that prohibit oxygen decline in sites associated with continual organic carbon inputs making them more resilient against climate-induced reductions in hypolimnetic oxygen. Consequently, although contributing less to the total carbon release from northern lakes when compared to polyhumic lakes, the low-DOC oligohumic lakes of the northern treeline are generally more prone to climate change impacts with increased potential of higher CH<sub>4</sub> emissions under reduced bottom water oxygen levels.

300

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302

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307

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517 **8 Tables**

518 **Table 1** Mean values and range (in brackets) in present-day wetland coverage and dissolved organic  
 519 carbon (DOC) in the 30 subarctic lakes in northeastern Finnish Lapland in comparison with Bray-  
 520 Curtis similarity and changes in N2 diversity and hypolimnetic oxygen between the surface (top)  
 521 and reference (bottom) samples.

	Pine and birch forest	Mountain birch woodland	Tundra
Wetland area of catchment (%)	5.8 (0...16.6)	15.6 (0.1...49.3)	2.9 (0...8.9)
DOC (mg l <sup>-1</sup> )	4.8 (2.2...9.6)	7.6 (2.5...16.6)	3.4 (1.7...7.4)
Similarity (Bray-Curtis)	0.50 (0.37...0.68)	0.55 (0.40...0.71)	0.44 (0.27...0.58)
Effective diversity change (N2)	0.12 (-5.03...2.35)	-1.69 (-7.64...6.54)	-0.74 (-8.61...7.59)
Oxygen change (mg l <sup>-1</sup> )	-3.0 (-7.0...4.0)	-1.9 (-7.0...0.5)	-4.5 (-9.6...-0.2)

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537 **9 Figure captions**

538

539 **Fig.1** Map. Study sites along the northern forest zones in northeastern Finnish Lapland. The  
540 numbers present lake codes, which are the same as in Rantala et al. (2016a).

541

542 **Fig. 2** Community compositions. Chironomid assemblages in the surface (top = gray) and reference  
543 (bottom = black) sediment samples in the 30 lakes in northeastern Finnish Lapland. The lakes are  
544 arranged according to their modern dissolved organic carbon (DOC) concentration and the  
545 threshold of 4 mg l<sup>-1</sup> (between oligo- and mesohumic) is marked with a dashed line.

546

547 **Fig. 3** Ecological changes. Bray-Curtis similarity (0 = dissimilar, 1 = similar) between the surface  
548 and reference chironomid assemblages in lakes in northeastern Finnish Lapland compared with  
549 Hill's (1973) N2 effective number of occurrences (diversity) and chironomid-inferred hypolimnetic  
550 oxygen along the dissolved organic carbon (DOC) gradient in the surface (top = gray) and reference  
551 (bottom = black) sediment samples.

552

553 **Fig. 4** Oxygen changes. Difference between chironomid-inferred hypolimnetic oxygen in the  
554 surface and reference sediment samples (negative values refer to decreased oxygen towards the  
555 present) compared with modern measured dissolved organic carbon (DOC) in the 30 lakes in  
556 northeastern Finnish Lapland. The samples are colored according to DOC thresholds of 4 and 6 mg  
557 l<sup>-1</sup> (separating oligo-, meso- and polyhumic sites) and the site with anomalously high pH (an outlier)  
558 is marked with an x.